

"SPURS" in the North Atlantic Salinity Maximum

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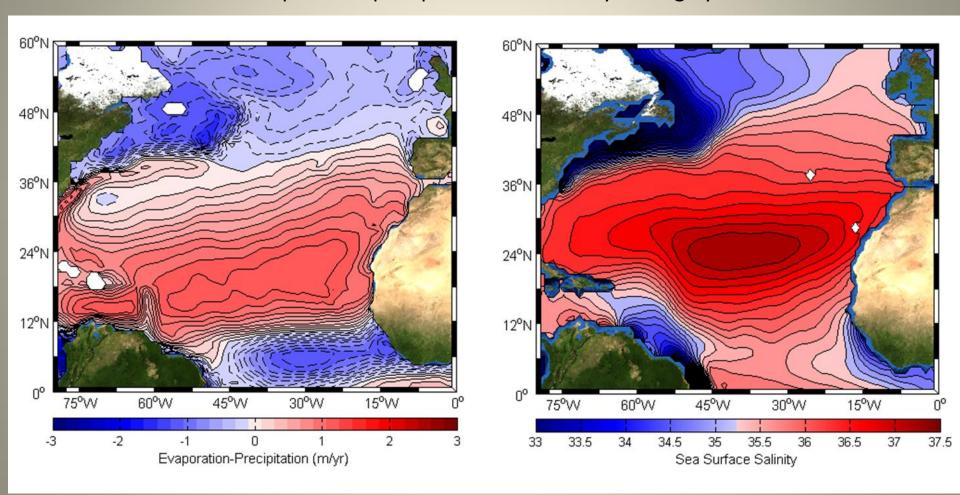








N. Atlantic evaporation-precipitation and salinity are highly correlated.



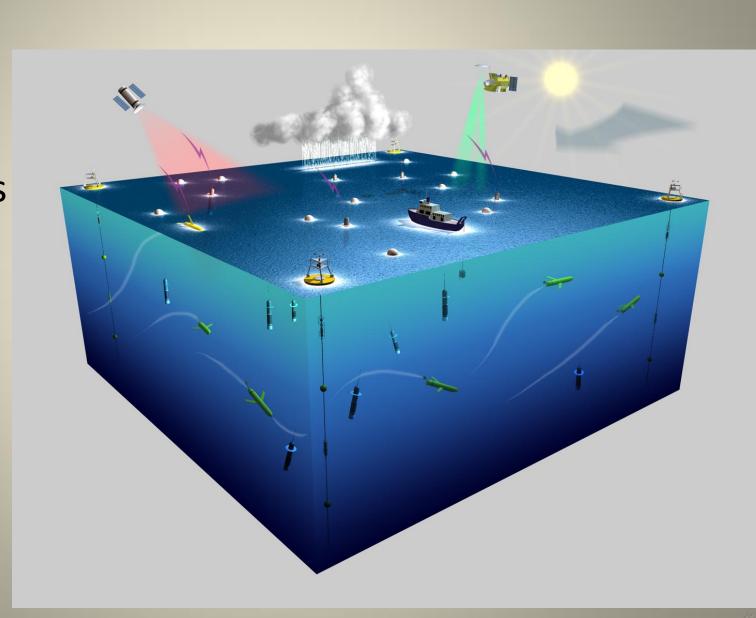
Note: the E-P zero line is close to vegetation/dry land boundary in Africa

Salinity Processes in the Upper ocean Regional Study

- Multi-agency (NASA, NOAA, NSF)
- Multi-national (US, Spain, France, Ireland)
- Designed to exploit new autonomous technology (Floats, gliders, wavegliders, drifters etc.)
- Linked to but not dependent on Aquarius and SMOS salinity satellites
- Cruises:
- French -Thalassa 8/16/12 9/13/12, Canaries Azores
- US Knorr 9/6/12-10/9/12, Woods Hole Azores (asset deployment)
- **US Endeavor 3/15/13 -4/15/13, Narragansett Narragansett (asset service)**
- Spanish Sarmiento de Gamboa 3/16/13-4/13/13, Canaries Azores
- **US Endeavor 9/19/13-10/16/13, Azores Narragansett (asset recovery)**

Challenge: Integrating a diverse set of measurements

- Floats
- Gliders
- Drifters
- Moorings
- Ships
- AUVs
- Satellites
- CTD &
 Micro structure
 profiling

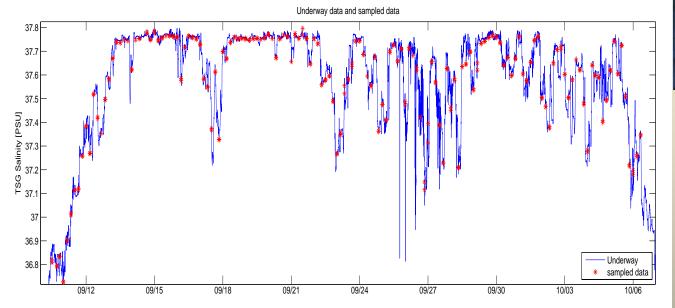


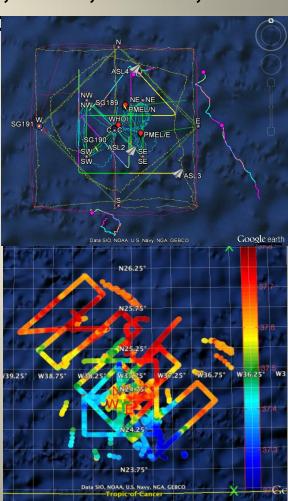
Sept.-Oct. cruise of R/V Knorr

•All assets deployed successfully: moorings, gliders, floats, drifters, Wave

Gliders, mixed layer float, microstructure profiler, etq

 Salinity max surveyed and found to have record high salinity ~37.8





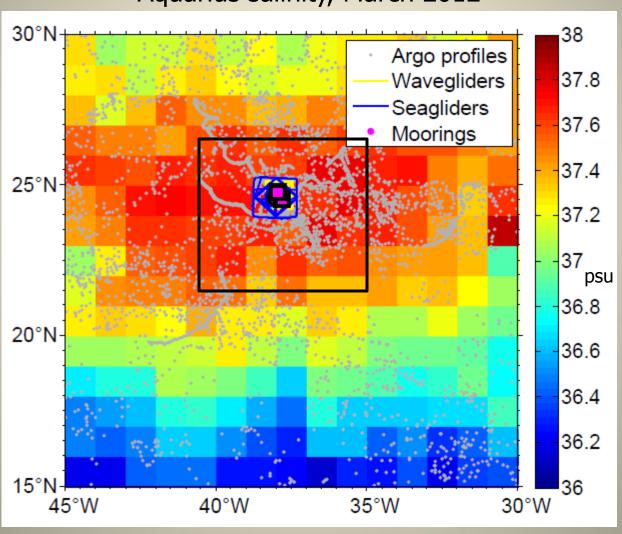
Two Approaches:

 1. Salinity budget at the central mooring on short time-scales (T. Farrar)

• 2. Average "isohaline" budget on long timescales (R. Schmitt)

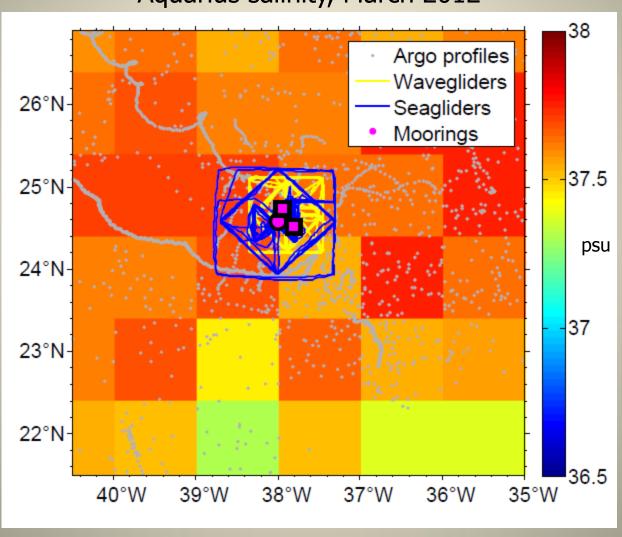
The SPURS "large box" (Tom Farrar)

Aquarius salinity, March 2012



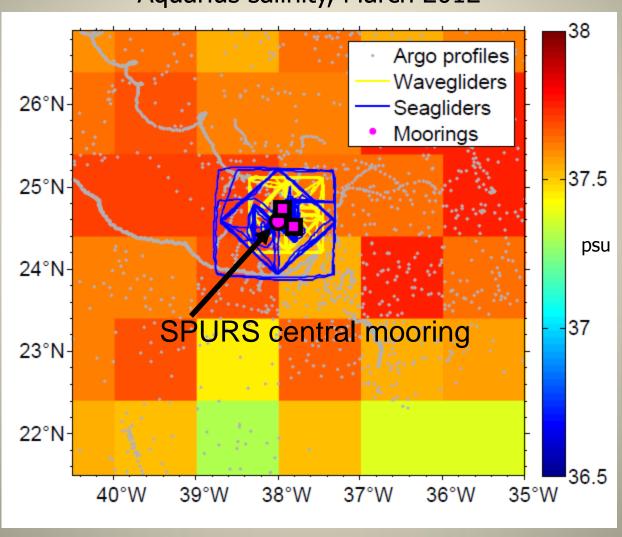
The SPURS "small box"

Aquarius salinity, March 2012

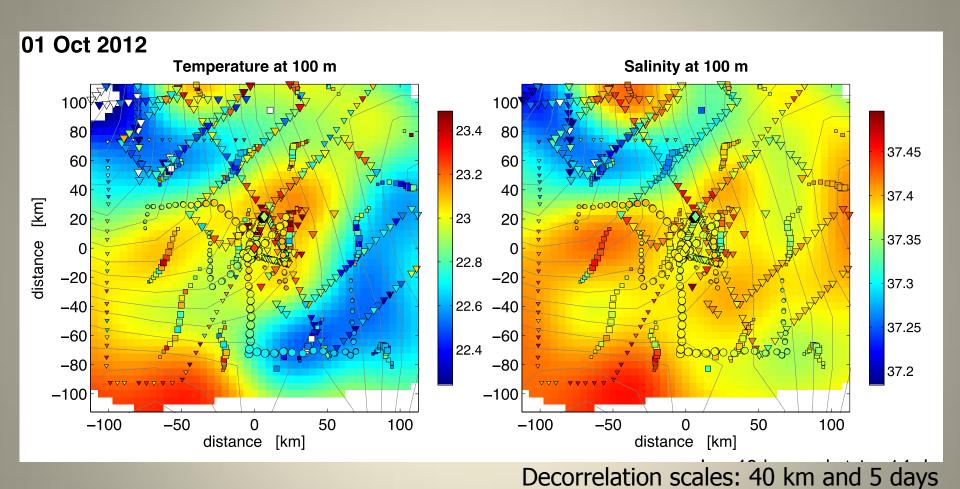


The SPURS "small box"

Aquarius salinity, March 2012



Gradient estimates, objective maps



Moorings

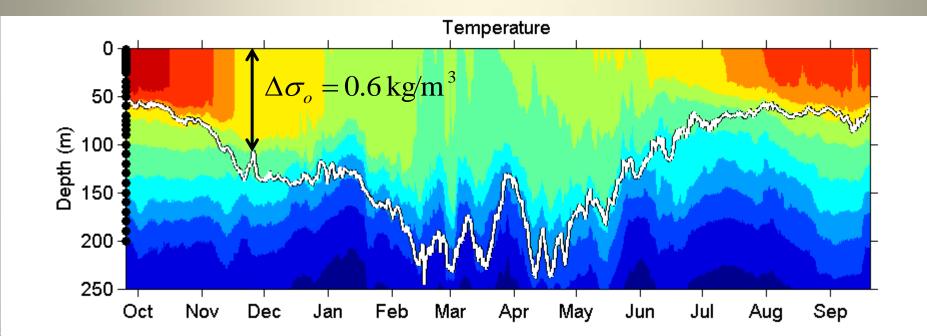
- Seagliders
- Argo floats
- **▽ UCTD**

Note: Size of data marker is scaled by when it was collected relative to the central time.

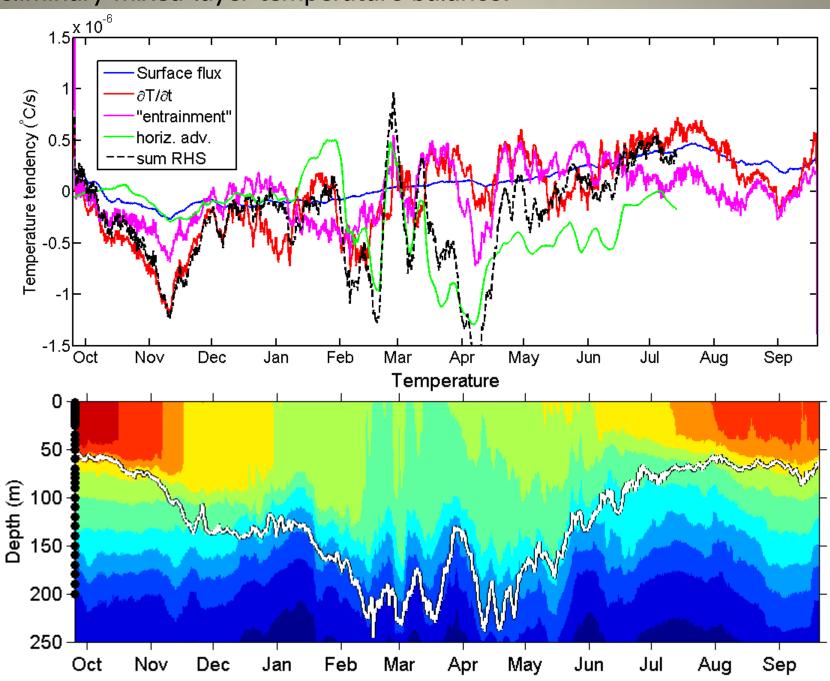
Preliminary mixed-layer temperature balance:

Evaluated for 3-week running average:

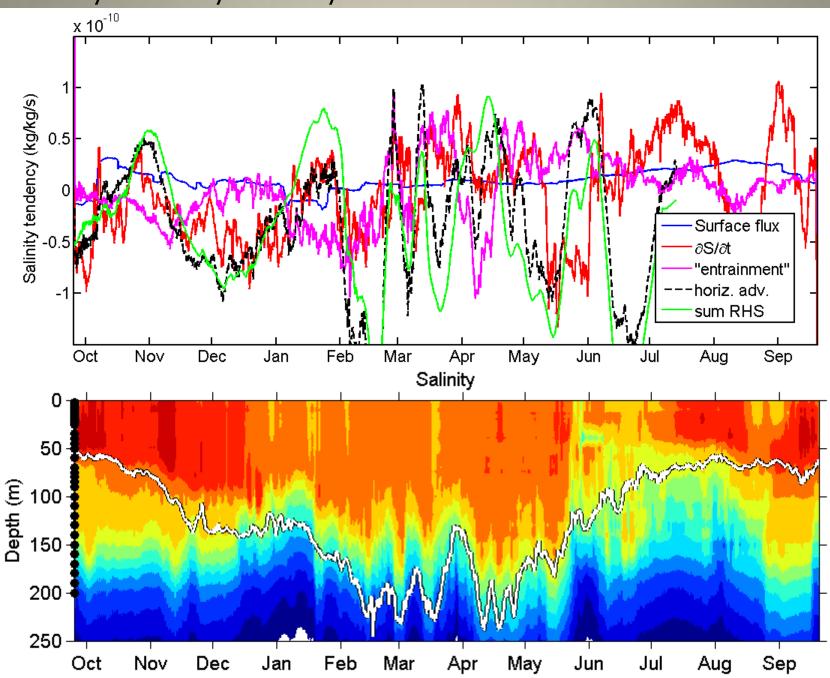
$$\frac{1}{\left(\frac{\partial \overline{T}}{\partial t}\right)} = \left(-\overline{\mathbf{u}} \cdot \nabla \overline{T}\right) + \left(\frac{\partial h}{\partial t}\right) w_{-h} + \mathbf{u}_{-h} \cdot \nabla h - \frac{Q_{-h}}{\rho c_{p} h} + \left(\frac{Q_{o}}{\rho c_{p} h}\right) + \frac{1}{h} \nabla \cdot \int_{-h}^{0} \hat{u} \hat{T} dz$$



Preliminary mixed-layer temperature balance:



Preliminary mixed-layer salinity balance:



Discussion of this preliminary budget:

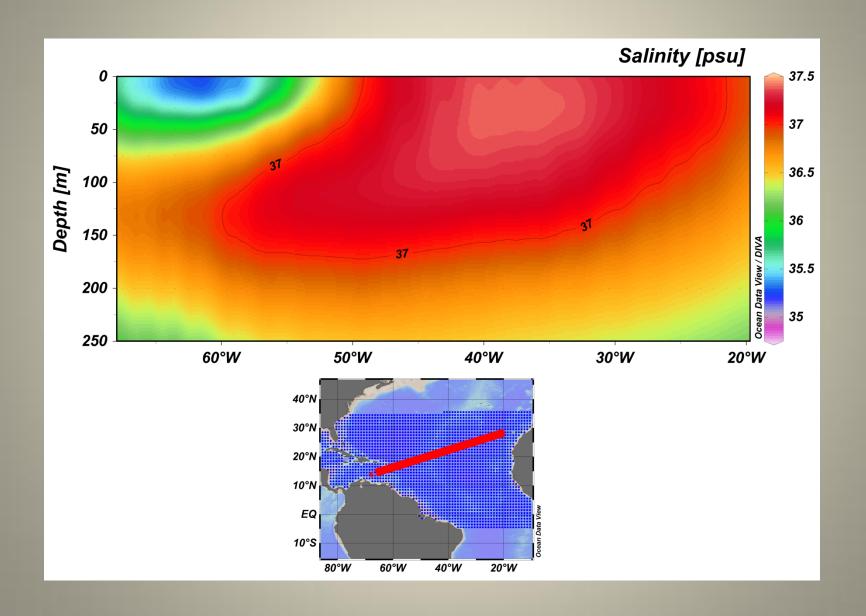
- (1) The mixed-layer temperature tendency appears to be roughly balanced by the sum of surface heat flux and mixed-layer "entrainment" terms
- (2) The mixed-layer salinity balance appears to be far from 1D on the 3-week timescales examined here
- (3) This difference is not too surprising because horizontal gradients of SSS are relatively strong. Surface heat fluxes will tend to damp SST anomalies, while surface freshwater fluxes are essentially independent of SSS.
- (4) The relative importance of terms in the budget equations will be different at different spatial and temporal scales— a larger goal of SPURS is to improve understanding of the various physical phenomena setting the mean and variability of SSS in the region.

Alternative Approach: An "isohaline" control volume

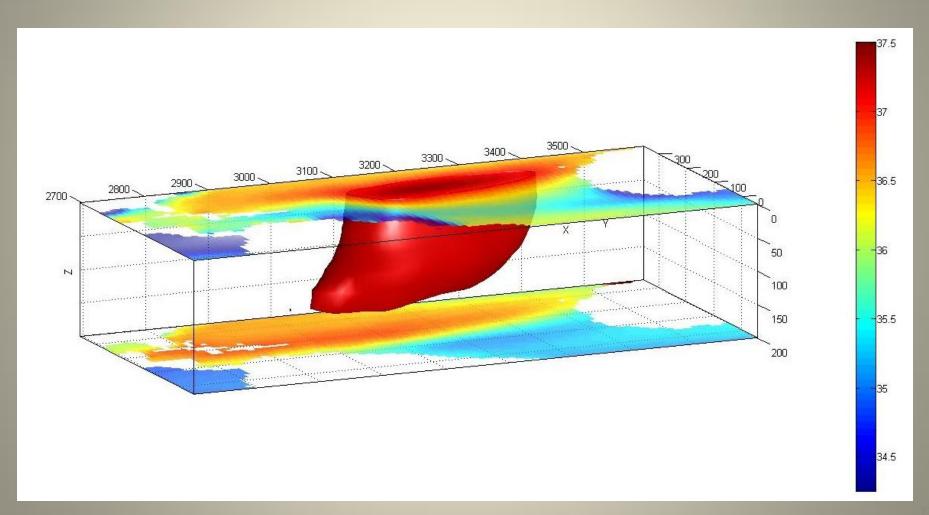
The idea goes back to Niiler and Stevenson (1982), "The heat budget of tropical ocean warm-water pools", *J. Mar. Res.*, *40*, 465–480.

Define a control volume by an isohaline, then the mean advective terms across this surface all have the same salinity and the balanced dynamical flows can be dropped from the equations (Ekman and Sverdrup circulations). Only the net water loss due to Evaporation minus Precipitation (E-P) counts in the mean advective salt balance. This must be compensated for by vertical and lateral mixing processes.

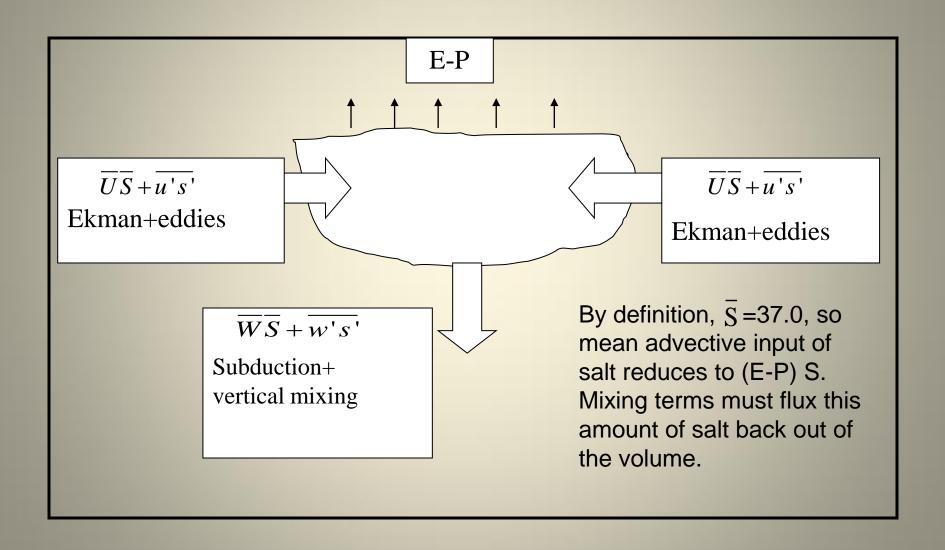
Vertical Structure of "Subtropical Underwater"



3-D Shape of the 37.0 Isohaline Salinity Maximum Characteristics



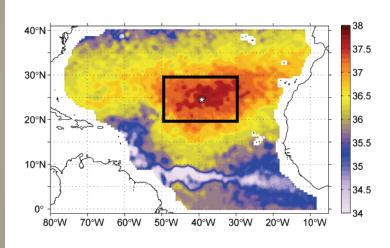
Salt Fluxes

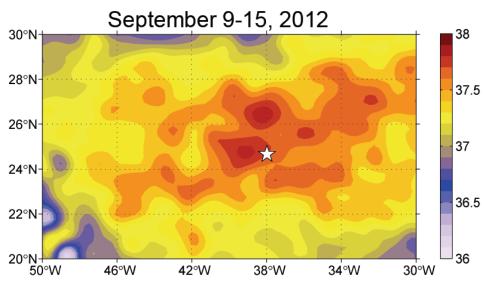


Motivation Eddy Activity clearly seen in SSS field

International Pacific Research Center, SOEST, University of Hawaii

Aquarius OI SSS for SPURS

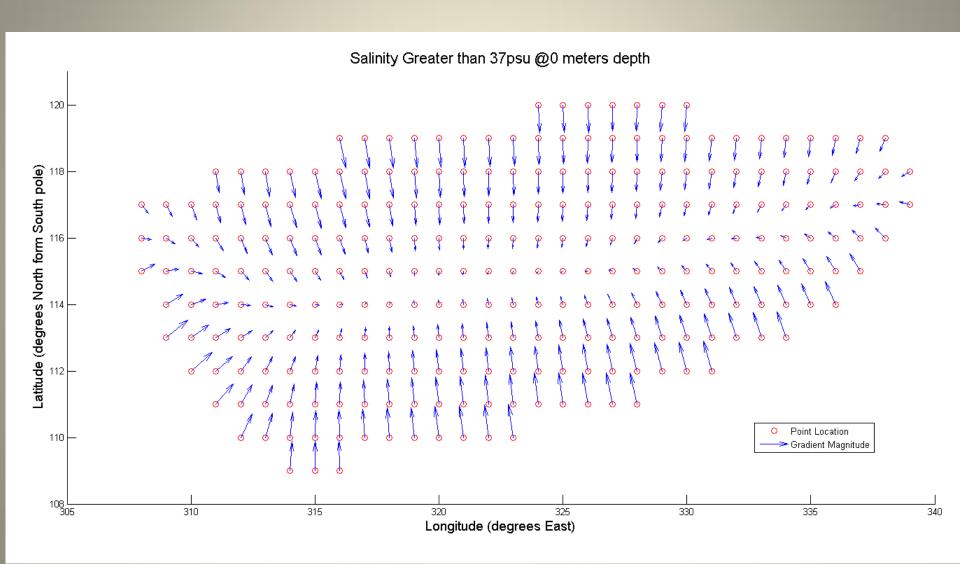




Data sources

- Use World Ocean Atlas 2009
- Define volume by S > 37.0
- Surface Area = $2.9 \times 10^6 \text{ km}^2$
- Net water loss from E-P= 0.11 Sv (Schanze, Schmitt and Yu, 2010)
- Net surface salt flux = 3.9 psu-Sv

Calculation of Ring Areas and Gradients



Diffusivity Calculations

- Salt Flux: $F_S = (E P) * S$
- · Balance:

$$\int F_S dA_S = \int K_V S_Z dA_B + \int K_H \nabla_2 S dA_W$$

Estimate K_V from microstructure measurements from VMP and T-Gliders, Estimate K_H from SSS drifters and models

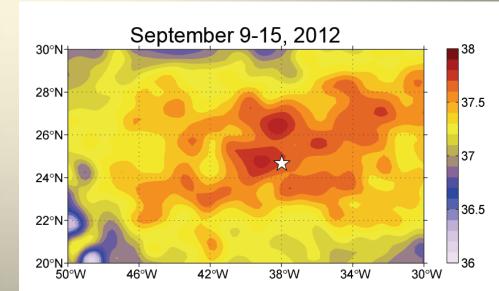
First results:

 Average Turbulent diffusivity from ~1000 T-Glider dives at the 37.0 isohaline yields:

$$K_{\rho} = 6.8 \times 10^{-5} \text{ m}^2/\text{s} \rightarrow \text{salt flux} = 1.05 \text{ PSU-Sv}$$

- Vertical mixing accounts for ~¼ of the salt flux put in at the surface.
- Requires $K_H \approx 5 \times 10^3 \,\mathrm{m^2 s^{-1}} \approx 0.1 \,\mathrm{ms^{-1}} \times 50 \,\mathrm{km}$

 $K_H \approx u'l$ (velocity scale × length scale)



Summary: SPURS-1 in the North Atlantic Salinity Maximum

- A well instrumented field program (with 5 cruises) was carried out in the NA Salinity Maximum, providing opportunities for two approaches to budget calculations.
- The moored time series, in combination with Lagrangian measures of the lateral gradients, allows point budgets to be evaluated on ~3 week time scales.
- The isohaline control volume is a promising approach for evaluating the large space and long time-scale balances in this unique oceanic closed control volume.
- Much analysis remains to be done!